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河西走廊地区民乐盆地的形成与改造: 对青藏高原东北缘白垩纪构造特征的启示

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Formation and destruction of the Minle Basin in the Hexi Corridor: insights into Cretaceous tectonic characteristics of the northeastern Qinghai-Tibetan Plateau

Abstract: [Objective] A series of sedimentary basins developed in the Hexi corridor along the northeastern margin of the Qinghai-Tibetan Plateau during the Early Cretaceous. These basins preserve critical records of pre-Cenozoic tectonic evolution in this region, forming the basis for understanding the growth mechanism of the Qinghai-Tibetan Plateau during the Cenozoic. However, the nature of these Early Cretaceous basins remains controversial, hindering consensus on late Mesozoic tectonics in this critical area. [Methods] This study focuses on the Minle Basin in the central Hexi Corridor and carried out a detailed field structural investigation in this area. Through structural analysis and paleo-stress reconstruction, the formation and destruction process of the Minle Basin was redefined. [Results] Our investigation reveals that Lower Cretaceous strata in the Minle Basin contain numerous syn-sedimentary normal faults that strike initially nearly north-south, indicating that the Minle Basin was a fault-bounded extensional basin controlled by nearly east-west extension during the Early Cretaceous. Subsequent deformation is recorded by NE-SW- and NW-SE-trending folds, coupled with shortening structures in Lower Cretaceous strata, indicating NE-SW- and NW-SE-trending horizontal shortening events during the Late Cretaceous. [Conclusion] The Early Cretaceous extension of the Minle Basin was consistent with the widespread extensional deformation and extensional stress direction across East

Asia during this period, which was a product of the retreat of the Paleo-Pacific Plate and the consequent mantle material flow. The bidirectional shortening during the Late Cretaceous resulted from the superimposition of the remote effects of simultaneous compression events that developed in the Tethys tectonic domain along the southern margin of the Eurasian Plate and the Pacific tectonic domain along the eastern margin of the Eurasian Plate. [Significance] The deformation mentioned above indicates that the influence of the subduction of the Paleo-Pacific Plate extended westward at least as far as the Hexi Corridor and the northern Qinghai-Tibetan Plateau.

Keywords: Northeastern Qinghai-Tibetan Plateau; Hexi Corridor; Minle Basin; Cretaceous; fault-bounded extensional basin

摘要: 青藏高原东北缘在早白垩世发育一系列沉积盆地, 这些盆地的构造特征记录了高原东北缘前新生代构造演化历史, 是认识青藏高原新生代生长机制的基础。但目前关于这些早白垩世盆地的性质目前仍存在诸多分歧, 直接影响了对该地区晚中生代构造特征的认识。以河西走廊中段民乐盆地为研究对象, 开展了系统的野外构造地质调查和古应力场恢复, 通过构造解析, 重新厘定了民乐盆地的形成和改造过程。研究发现, 民乐盆地下白垩统发育大量同沉积正断层, 其初始走向为近南北向, 指示民乐盆地在早白垩世为受控于近东西向伸展作用的断陷盆地。民乐盆地下白垩统呈现出北东—南西向和北西—南东向 2 组轴向的褶皱, 结合地层中发育的缩短构造, 揭示该盆地在晚白垩世遭受了北东—南西向和北西—南东向双向水平缩短作用。民乐盆地早白垩世的伸展作用, 与东亚地区早白垩世广泛发育的伸展变形及伸展应力方向具有一致性, 是古太平洋板块回撤以及由此引发的地幔物质流动的产物。晚白垩世遭受的双向缩短, 则是欧亚板块南缘特提斯构造域和东缘太平洋构造域同时期挤压事件远程效应叠加的结果。上述变形特征表明, 古太平洋构造域的影响范围至少可向西扩展至河西走廊及青藏高原北部地区。

关键词: 青藏高原东北缘; 河西走廊; 民乐盆地; 白垩纪; 伸展断陷盆地

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0 引言

青藏高原是现今地球上规模最大、海拔最高的高原, 其构造演化过程吸引着众多地质学家的关注 (Molnar and Tapponnier, 1975; Dewey et al., 1988; Molnar et al., 1993; Yin and Harrison, 2000; Tapponnier et al., 2001; Yin, 2006)。现有研究表明, 青藏高原的形成和扩展主要发生在新生代 (Tapponnier et al., 2001; Yin, 2006; Ding et al., 2022; Wang et al., 2022a; Xiong et al., 2022), 但前新生代构造被认为对高原及其以北地区的新生代构造变形有着重要的控制作用 (Dewey et al., 1988; Kong et al., 1997; Yin and Harrison, 2000; Yin et al., 2008; Cunningham et al., 2016; Tian et al., 2016; Cheng et al., 2019; Cunningham and Zhang, 2020; 张进等, 2023; An et al., 2025)。

中生代期间, 受特提斯构造域微陆块/岛弧持续向北俯冲拼贴于欧亚板块南缘的影响, 青藏高原地区构造变形主要由近南北向的缩短变形主导 (Dewey et al., 1988; Yin et al., 2008; Li et al., 2025; Li and Jepson, 2025; Li et al., 2026)。这些构造事件影响并控制了高原内部及周边地区盆地的发育, 形成了一系列陆内挤压相关盆地 (Hendrix et al., 1996; Ritts and Biffi, 2000; Cheng et al., 2019; Hu et al., 2022)。而在青藏高原以东的东亚地区, 中生代构造变形则表现为多期次缩短和伸展变形的交替发育 (董树文等, 2019; 张岳桥和董树文, 2019; Li et al., 2020, 2024; Suo et al., 2020; Zhang et al., 2020, 2021a, 2021b, 2022, 2025; 张北航等, 2021; Xu et al., 2025; 张进等, 2025; Feng et al., 2026)。尤其在早白垩世, 东亚地区发育大规模伸展变形, 形成了广泛分布的变质核杂岩、伸展断陷盆地等构造 (Meng et al., 2003; Davis and Darby, 2010; Wang et al., 2011; Lin and Wei, 2020; Zhang et al., 2021b, 2022; Liu et al., 2021; Zhang, 2026)。青藏高原东北缘是高原向外扩展的主要前缘地带, 该区

域发育有一系列早白垩世沉积盆地。以祁连造山带为界，其南侧的柴达木盆地在早白垩世为挤压拗陷盆地（Cheng et al., 2019; 张晨雨, 2020），而北侧的河西走廊和阿拉善地区，则在早白垩世发育大规模伸展盆地以及火山活动（137~99 Ma; 卫平生等, 2005; 钟福平等, 2011; Zhang et al., 2018, 2021b, 2022; 陈志鹏等, 2019; Hui et al., 2021; Zhang, 2026）。这2种截然不同的构造环境是否同时存在？若存在，二者如何相互影响？其影响边界分布于何处？控制2类盆地发育的构造机制又是什么？这亟需回答的科学问题是理解青藏高原及周边地区前新生代构造演化的关键，也是厘清东亚地区中生代构造格局的重要内容。

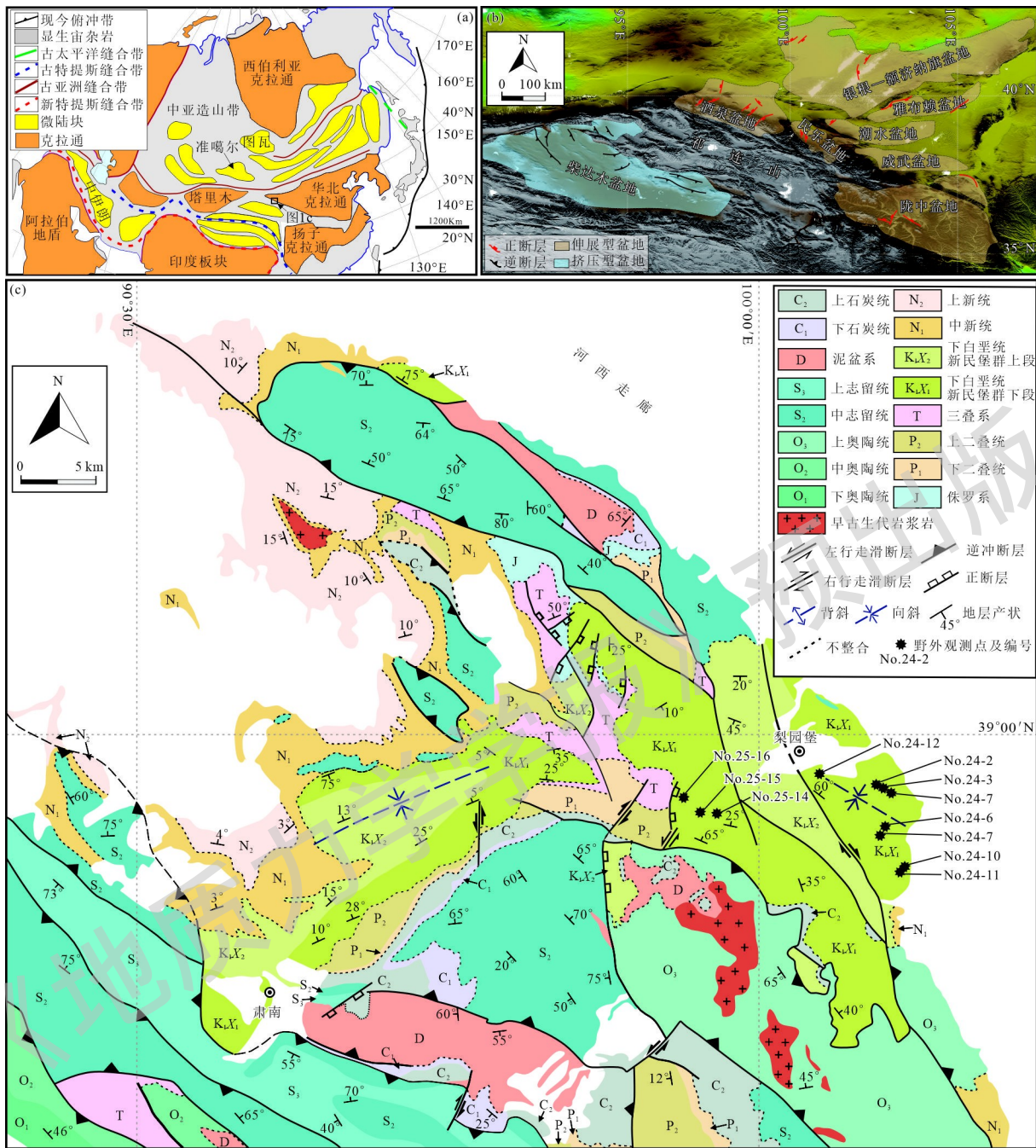
祁连山作为青藏高原的北侧边界，处于上述早白垩世2种构造体制的过渡区域。紧邻祁连山北侧的河西走廊地区发育一系列早白垩世沉积盆地（如民乐盆地、酒泉盆地等），明确这些沉积盆地的性质，有助于限定该地区在早白垩世是否存在2种不同构造环境及其成因机制。然而，目前对这些盆地性质仍存在不同认识：部分学者认为这些盆地是挤压环境下的前陆盆地（Wang and Coward, 1993; 陈宣华等, 2019a; 邵浩浩等, 2019; Wang et al., 2022b）；另有学者认为这些盆地是挤压环境下形成的走滑拉分盆地（王同和, 1987; Vincent and Allen, 1999; Yang et al., 2025）；还有部分学者认为这些盆地形成于伸展环境，属于伸展断陷盆地（李奋其, 2003; 王崇孝等, 2005; 王晓丰等, 2008; Zhang et al., 2018）。不同观点的存在，制约了对该地区早白垩世构造环境的深入认识。

针对上述问题，以河西走廊中部的民乐盆地为研究对象，开展了详细的野外构造地质调查，识别出大量同沉积构造及后期改造变形。通过构造解析，结合古应力场反演，恢复了盆地的沉积和改造过程；同时探讨了相应的构造背景，为青藏高原东北缘地区晚中生代构造演化提出了新的认识。

1 地质背景

1.1 大地构造位置及构造演化概况

民乐盆地位于青藏高原东北缘、河西走廊中段（图 1a、1b），南北两侧分别以北祁连逆冲断裂和龙首山南缘断裂为界，东西两端分别以永固-大黄山隆起和榆木山隆起为界，平面上整体呈菱形展布。盆地所在的河西走廊为现今青藏高原东北缘扩展前缘，其南侧的北祁连造山带为早古生代祁连洋（原特提斯洋）俯冲闭合的产物（Song et al., 2013），并在中一晚三叠世和早白垩世先后经历了缩短和伸展构造事件（Chen et al., 2003; 陈宣华等, 2019a, 2019b），民乐盆地主要形成于这个阶段。新生代，随着印度-欧亚板块的碰撞导致青藏高原的形成和扩展，北祁连造山带再次复活，发生强烈的挤压逆冲变形，在河西走廊南缘形成北祁连逆冲断裂（Yin and Harrison, 2000; Chen et al., 2003; Wang et al., 2016, 2020; 吴晨等, 2023; Wu et al., 2025）。



a—亚洲大地构造简图；b—青藏高原东北缘早白垩世盆地类型及分布；c—民乐盆地区域地质图

图1 亚洲大地构造简图、青藏高原东北缘早白垩世盆地类型、分布及民乐盆地区域地质图

Fig. 1 Sketched tectonic map of Asia, style and distribution of the early Cretaceous basins around the northeastern Qinghai-Tibetan Plateau and regional geologic map of the Minle Basin

(a) Sketched tectonic map of Asia; (b) Style and distribution of the early Cretaceous basins around the northeastern Qinghai-Tibetan Plateau; (c) Regional geologic map of the Minle Basin

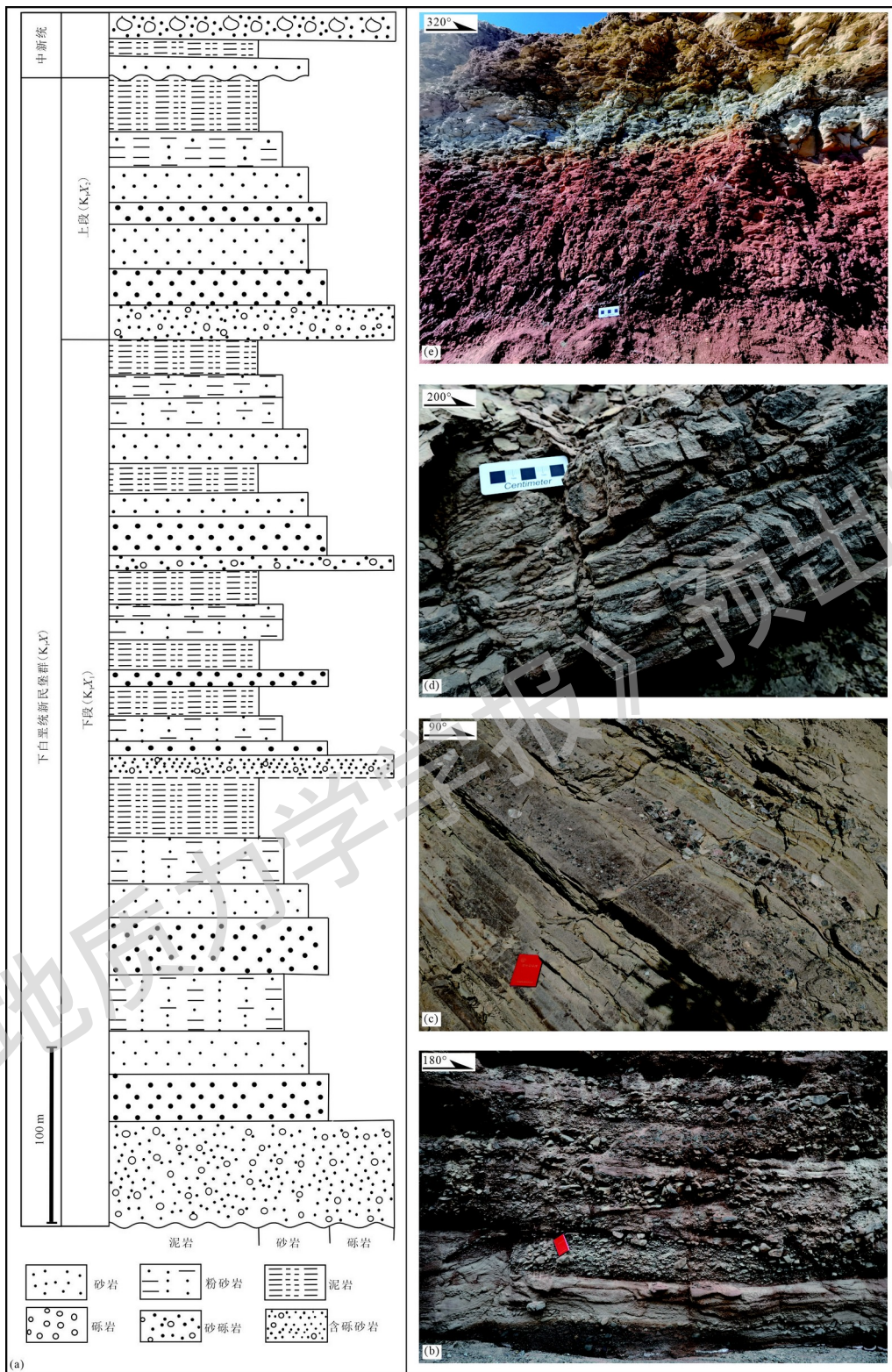
1.2 地层概况

民乐盆地内下白垩统为新民堡群（图1c），主要由砾岩、砂岩、粉砂岩和泥岩组成（甘肃省地质调查院，2021）。根据岩性，可分为上、下2段（图2a）。新民堡群下段（K₁X₁）下部以厚层状砾岩为主，局部夹岩屑砂岩层；砾岩为颗粒支撑，砾石磨圆差（图2b）；岩屑砂岩为粗粒结构，碎屑呈次棱角状或次磨圆状，局部发育交错层理。K₁X₁上部主要由砂岩和粉砂岩组成（图2c），整体呈现向上变细的趋势。新民堡群上段

(K_1X_2) 相较于下段, 细粒沉积增多。 K_1X_2 下部岩性主要为砂岩, 并呈现出多个粒度向上变细或向上变粗的旋回; 向上地层中一细砂岩和粉砂岩逐渐增多 (图 2d)。 K_1X_2 上部岩性主要为 200 m 厚的杂色互层状粉砂岩和泥岩, 地层呈现出浅灰色、黄绿色和暗红色 (图 2e)。最上部地层再次以厚层含砾粗砂岩和砂岩为主 (图 2a)。

民乐盆地缺失上白垩统, 中新统角度不整合覆盖于下白垩统之上, 主要由含砾砂岩、砂岩、砂质泥岩和泥岩组成 (图 2a); 上新统岩性为互层状粉砂岩和砾岩; 第四系主要为冲洪积沉积和风成黄土 (甘肃省地质调查院, 2021)。

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a—民乐盆地下白垩统地层柱状图；b—厚层砾岩；c—厚层粗砂岩；d—中薄层粉砂岩；e—杂色泥岩
图2 民乐盆地下白垩统新民堡群地层柱状图及主要岩性野外露头特征

Fig. 2 Stratigraphic Column of the Lower Cretaceous Xinminpu Group and Typical Lithologies in the field

(a) Stratigraphic column of the Lower Cretaceous Xinminpu Group in the Minle Basin; (b) Thick-bedded conglomerate; (c) Thick-bedded coarse sandstones; (d) Medium-thin-bedded siltstone; (e) variegated mudstone

2 研究方法

此次研究中,在开展野外工作前,通过对张掖梨园堡及周边地区卫星影像进行解译,识别并确定地层形迹和褶皱变形特征。在野外构造地质调查过程中,在每个构造观测点系统测量并收集相关的构造要素数据,包括地层产状、断层面产状和擦痕产状。断层运动学性质主要根据断层面及断层带内发育的运动学指示标志来判断,包括擦抹晶体、里德尔剪切破裂以及断层两盘标志层的错断等。古应力场恢复主要通过断层面滑动矢量数据、褶皱枢纽产状等构造要素数据来确定。

此次研究中,通过 Stereonet 11 (Allmendinger et al., 2012) 软件来实现构造数据的可视化;利用 FaultKin 8.1 (Allmendinger et al., 2012) 软件对断层面滑动矢量数据进行计算,实现古应力场数据的恢复。

3 构造变形特征

民乐盆地下白垩统记录了不同时期的变形,目前填图尺度的变形表现为北西—南东向褶皱和北东—南西向褶皱(图 1c),2 个方向的褶皱在梨园堡附近形成了叠加(东部被第四系覆盖)。除了上述的大型变形外,露头尺度也发育多种类型的构造变形,如生长断层、软沉积变形以及后期叠加于其上的多阶段的逆冲和走滑断层。

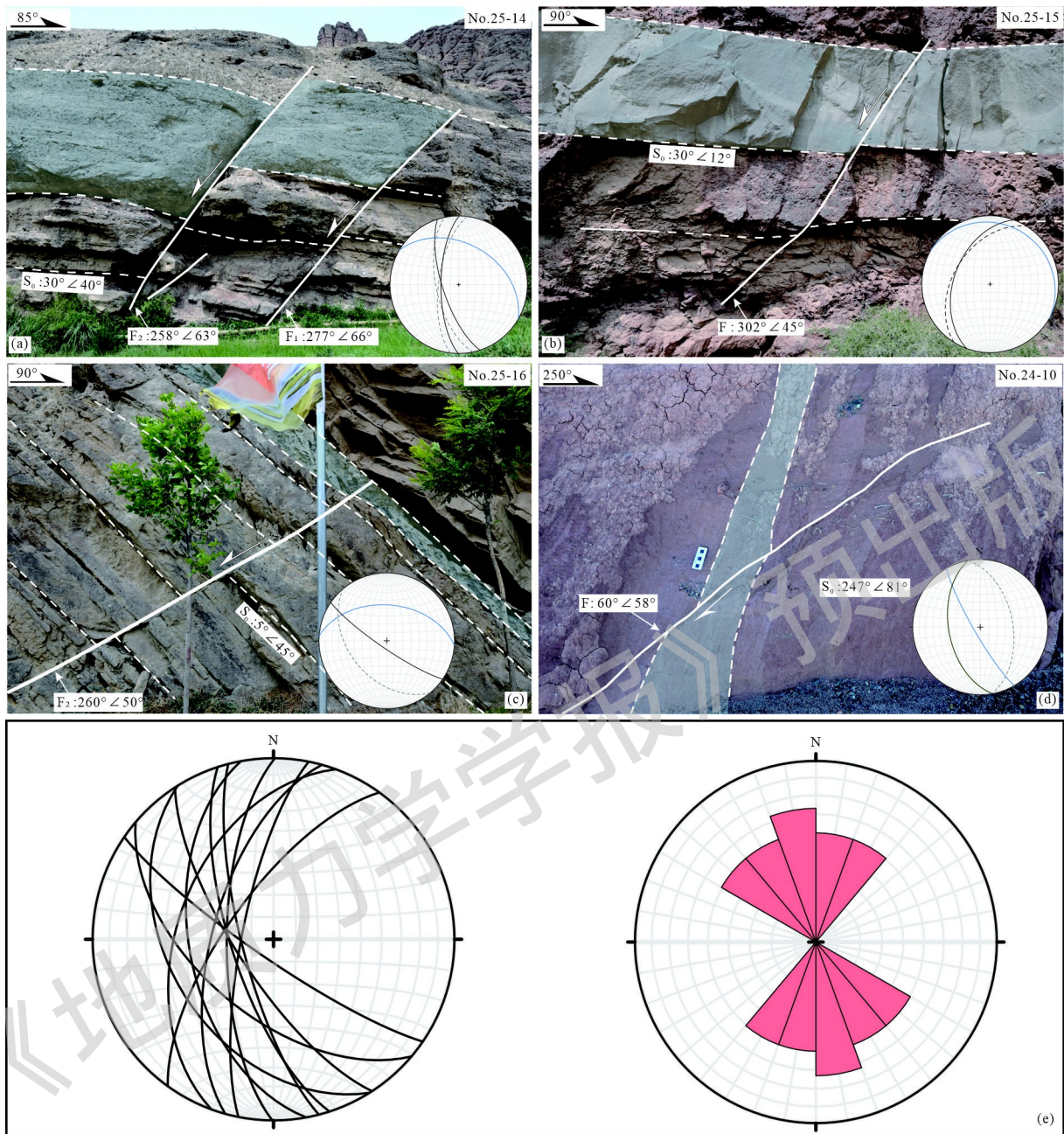
3.1 生长断层

民乐盆地下白垩统新民堡群下部(K_1X_1)岩性以厚层砾岩和含砾粗砂岩为主,在最下部的砾岩层和含砾粗砂岩层中,发育多条与地层大角度相交的断层,断层两侧地层大多表现出拖曳褶皱样式,指示断层为正断性质。在 No.25-14 和 No.25-15 观测点,由于后期剥蚀,断层顶部未完整出露(图 3a、3b)。野外可见断层切割整个露头地层,同时断层上盘的砾岩、砂岩或含砾粗砂岩地层厚度明显大于下盘对应层,指示断层活动可能控制着地层的沉积(图 3a、3b)。在 No.25-16 观测点,断层露头较好,完整揭露出整个断层结构。该观测点处断层切割了多层含砾粗砂岩,沿断层面由下往上,断距逐渐减小,断层顶部终止于一层黑色页岩中,黑色页岩在断层面两侧表现出一定的厚度变化,页岩顶部被一厚层粗砂岩覆盖,砂岩底面完整平直,未被断层错断(图 3c)。这一构造特征表明,该断层活动可能与最上部的地层沉积同期,并在黑色页岩沉积完成后停止活动。

在下白垩统新民堡群上部(K_1X_2)杂色砂岩与泥岩互层地层中,同样也发育有生长断层。在 No.24-10 观测点,岩性主要为中层状紫红色泥岩与砂岩互层,地层陡立且发生倒转。地层中发育一条现今露头表现为逆冲性质的低角度断层(图 3d)。断层两盘对应层厚度存在差异,现今下盘紫红色泥岩厚于上盘对应层。根据露头处地层倒转的特征,将地层恢复为水平状态时,该逆冲断层则为正断层,且恢复后的上盘地层厚度大于下盘,同样表现为同沉积正断层特征。

野外测量结果显示,这些同沉积正断层产状变化较大:在地层近水平地区,这些断层表现为高角度特征;在地层倾角较大或地层陡立倒转地区,断层表现为低角度特征(图 3)。根据断层发育位置的地层产状,对这些断层产状进行恢复后发现,当地层恢复至初始水平状态时,这些断层的产状基本一致,呈现出近南北走向(图 3e)。这一产状特征表明,这些同沉积正断层形成于近东西向水平伸展作用下。

需要说明的是,生长断层在民乐盆地下白垩统不同层位广泛发育,如在冰沟丹霞地区的下白垩统下部厚层砾岩中以及在梨园堡地区的中上部地层中均有发育,这说明早白垩世时期,伸展作用控制着民乐盆地的形成,且伸展方向为近东西向,与早白垩世整个东亚区域性的伸展方向(北西—南东向或北西西—南东东向)相近(张进等, 2025)。



图中所有赤平投影为下半球投影；蓝色实线代表地层产状，灰色虚线代表实测生长断层产状，黑色实线表示旋转至初始状态的生长断层产状

a—生长断层两盘同沉积砾岩厚度发生变化；b—生长断层两盘同沉积粗砂岩厚度发生变化；c—生长断层两盘同沉积页岩厚度发生变化，顶部被厚层砂岩覆盖；d—倒转地层中发育的生长断层现今表现为逆断层性质；e—赤平投影和玫瑰花图展示生长断层初始产状特征

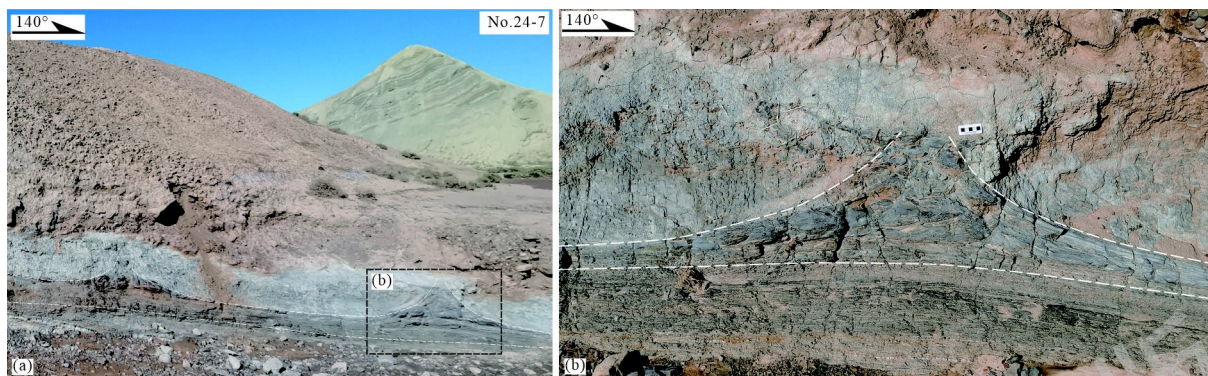
图3 民乐盆地下白垩统中生长断层构造特征（野外观测点位置见图1c）

Fig. 3 Structural Characteristics of Growth Fault in the Lower Cretaceous Strata of the Minle Basin (See Figure 1c for the field observation sites locations. Lower hemisphere, equal area projections apply for all stereoplots in this paper. The blue solid lines represent the bedding, the gray dashed lines and the black solid lines represent the measured and initial attitude of the growing fault, respectively)

(a) Syn-rift conglomerate show thickness variations across growth fault; (b) Syn-rift coarse sandstone show thickness variations across growth fault; (c) Syn-rift black shale show thickness variations across growth fault and is covered by thick bedded sandstone; (d) The growth faults developed in the overturned strata now express as reverse faults; (e) Stereo-projection and rose diagram of initial attitude of

the growth faults

除上述与地层沉积同期发育的生长断层外，下白垩统中还发育一系列软沉积变形，如“帐篷”构造等（图4），这些变形与沉积同时，也属于同沉积变形，反映沉积过程中存在构造扰动，但不能体现区域性构造的影响。



a—民乐盆地下白垩统软沉积变形发育层位；b—灰黑色粉砂质泥岩中发育的“帐篷”构造

图4 民乐盆地下白垩统软沉积变形（观测点位置见图1c）

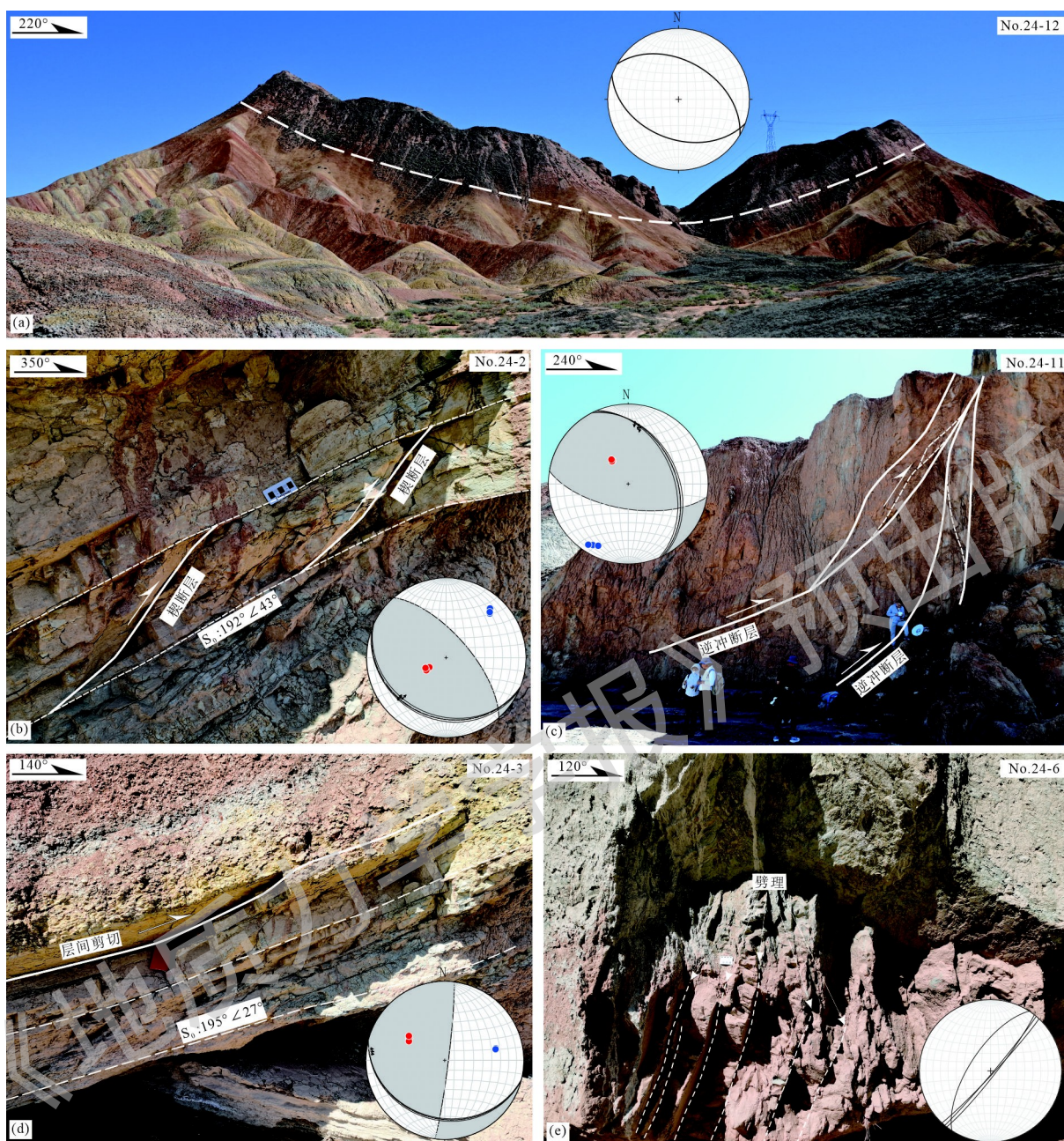
Fig. 4 Soft-deformation developed in the Lower Cretaceous Strata in the Minle Basin (See Figure 1c for the field observation site location)

(a) Position of soft-deformation developed in the Lower Cretaceous; (b) Tepee-like structures developed in grayish-black sandy mudstone

3.2 后期变形

民乐盆地下白垩统在肃南县北部以及梨园堡地区呈现出大型的褶皱（图5a），表明下白垩统在沉积之后经历了缩短变形。此次研究在下白垩统中发现了多种类型的缩短构造（图5）。

野外多处露头可见下白垩统厚层砂岩中发育一组与层面小角度相交逆冲断层，该组断层大多局限于某一沉积层中，断层面与地层顶底面逐渐变缓，被切割的地层沿断层发生一定位移（图5b）。这些断层多为地层在水平状态下遭受水平缩短作用时形成的楔断层，反映了盆地在沉积结束后经历的初始缩短变形（Weil and Yonkee, 2023）。根据断层面滑动矢量数据的反演结果，这些楔断层形成于北东—南西向水平挤压应力作用下（图5b）。除楔断层外，下白垩统中还发育一系列高角度逆冲断层，主要发育于靠近褶皱核部位置（图5c），是地层持续褶皱过程中的调节断层（breakthrough fault）。根据断层面滑动矢量数据反演得到这些高角度逆冲断层分别形成于北东—南西向和近东西向水平挤压应力作用下（图5c）。此外，野外露头还可见沿地层层面发生的层间滑动，为地层褶皱过程中层间剪切的产物。根据层间滑动形成的擦痕，反演得到近东西向水平挤压应力方向（图5d）。除逆冲断层外，野外露头还可见下白垩统厚层泥岩中发育一组劈理，走向北东—南西向，反映了北西—南东向水平挤压作用（图5e），这与褶皱的形成可能存在关联。



图中所有断层机制解中黑色实线代表断层面，蓝点代表P轴，红点代表T轴

a—梨园堡地区北西走向向斜转折端；b—楔断层；c—褶皱核部逆冲断层；d—细砂岩层之间发育的层间剪切；e—泥质粉砂岩中发育的北东—南西走向密集劈理

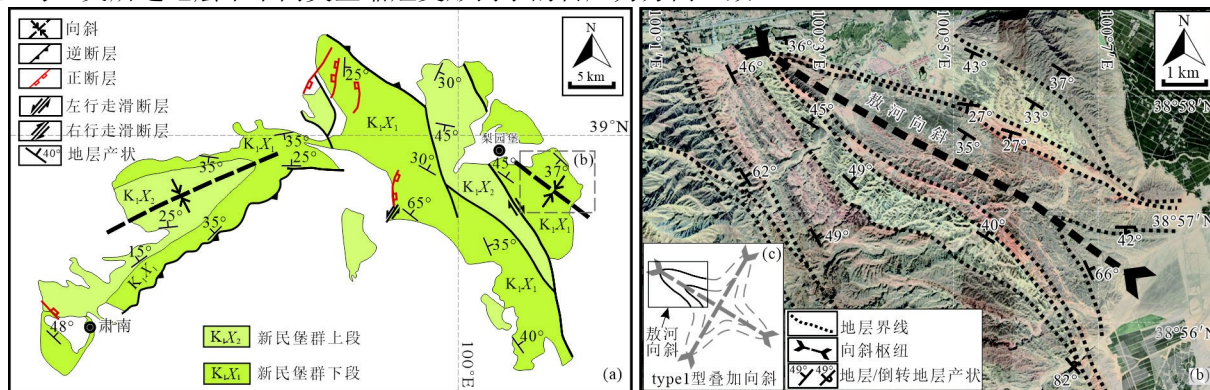
图5 民乐盆地下白垩统中的缩短变形（观测点位置见图1c）

Fig. 5 Shortening Deformation of the Lower Cretaceous strata in the Minle Basin (See Figure 1c for the observation sites locations. In the fault plane solutions, black solid lines represent fault planes, blue dots represent P-axis, and red dots represent T-axis)

(a) Hinge zone of NW-trending syncline in the Liyuanpu Area; (b) Wedge fault; (c) Reverse faults developed in the fold core zone; (d) Slip along fine sandstone layers interfaces; (e) NE-SW-strike intensive cleavage developed in argillaceous siltstones

此外，民乐盆地下白垩统发育明显褶皱变形。在梨园堡地区，下白垩统表现为北西—南东走向、向南东方向倾伏的向斜（敖河向斜；图6a、6b）。且该向斜在平面上表现为向南东方向开口的喇叭状（图6b），向斜两翼地层产状沿褶皱枢纽走向逐渐变化：由北西向南东方向，两翼地层走向逐渐偏离枢纽，分别向北东和南西方向转向（图6b）。受后期剥蚀和第四系覆盖的影响，敖河向斜继续向南西方向的延伸未完全出露，但是目前残留的形态特征揭示了一个不完整的叠加褶皱的形态（图6b），是北东—南西向和北西—南东向缩短作用

叠加的产物，为 type1 型叠加褶皱（Ramsay and Huber, 1987）。同时，肃南县北侧下白垩统表现为枢纽北东—南西走向、向南西倾伏的向斜（图 6a），与梨园堡地区北西—南东走向的向斜近直交。这 2 处向斜枢纽的产状与敖河向斜揭示出的叠加向斜枢纽产状一致，均反映了北东—南西向和北西—南东向 2 个近直交方向的缩短作用，与上文所述地层中不同类型缩短变形揭示的古应力方向一致。

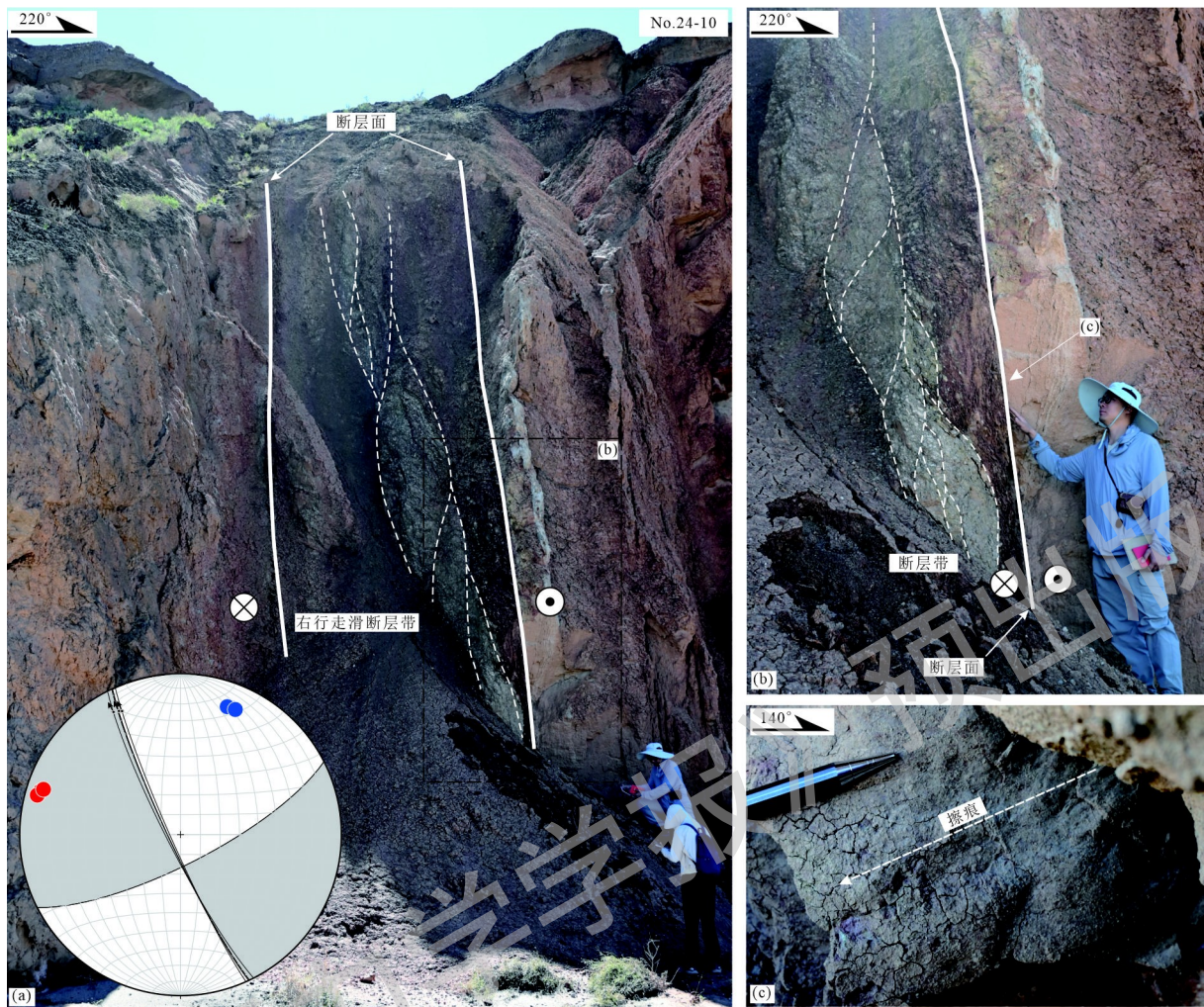


a—下白垩统形成轴向北西—南东和北东—南西近直交的 2 组褶皱；b—梨园堡地区叠加向斜；c—type1 型叠加向斜样式
图 6 民乐盆地下白垩统褶皱变形特征

Fig. 6 Fold patterns of the Lower Cretaceous in the Minle Basin

(a) The Lower Cretaceous formed two sets of nearly orthogonal folds with axial trends of NW-SE and NE-SW respectively; (b) Superimposed syncline in the Liyuanpu area; (c) type 1 fold interference pattern

除逆冲断层外，下白垩统中还发育一系列走滑断层。这些走滑断层大多为北西—南东走向，与北西—南东轴向褶皱枢纽近平行发育，切割下白垩统；在局部地层陡立地区，断层面与地层层面近平行（图 7）。部分露头可见走滑断层形成宽约 3 m 的断层破碎带，带内劈理发育（图 7）。研究区内的走滑断层主体为右行走滑，个别为左行走滑，兼具一定的逆冲分量，同样形成于挤压环境。部分露头测得的断层数据指示北东—南西向水平挤压应力（图 7a），与上文所述缩短变形的古应力场方向一致。这些走滑断层可能与上述缩短构造为同时期变形，是褶皱持续变形晚期的调节变形，但也不排除其为更晚期变形的可能性，目前尚无定论。



a—断层带及断层机制解（断层机制解中黑色实线代表断层面，蓝点代表P轴，红点代表T轴）；b—断层带内劈理发育；c—断层面擦痕指示右行走滑

图7 民乐盆地下白垩统中北西走向右行走滑断层（观测点位置见图1c）

Fig. 7 NW-trending dextral strike-slip fault developed in the Lower Cretaceous in the Minle Basin (See Figure 1c for the observation site location).

(a) Fault zone and fault plane solution (In the fault plane solutions, black solid lines represent fault planes, blue dots represent P-axis, and red dots represent T-axis); (b) Cleavage developed in the fault zone; (c) Slickenlines on the fault plane indicate dextral strike-slip

4 讨论

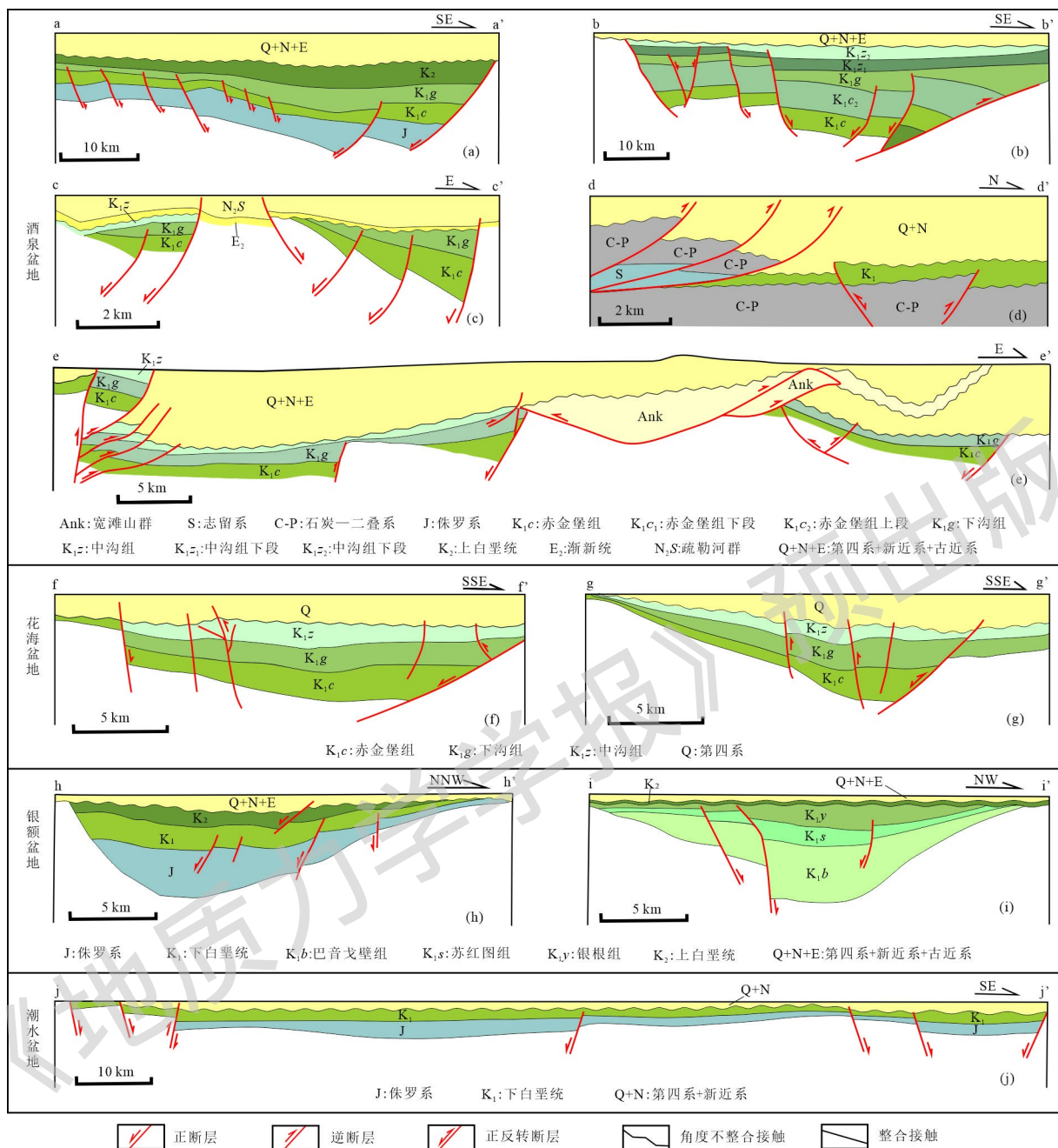
4.1 民乐盆地的形成与改造

4.1.1 早白垩世构造事件

相关学者（Wang and Coward, 1993; 邵浩浩等, 2019）研究认为，祁连山以北的河西走廊地区在早白垩世处于挤压环境，形成了一系列前陆盆地，如酒泉盆地、民乐盆地和平山湖盆地等。该期挤压在民乐盆地西侧还形成了榆木山飞来峰（陈宣华等, 2019b）。另有部分学者（王同和, 1987; Vincent and Allen, 1999）指出，河西走廊地区的早白垩世盆地为右行走滑拉分盆地，主要应力环境为北东—南西向挤压（Vincent and Allen, 1999）。最近，Yang et al. (2025) 根据断层运动学和年代学的研究，提出河西走廊地区在早白垩世发育由左行走滑断层控制的转换挤压系统，该系统中局部拉伸作用形成了包括民乐盆地在内的早白垩世沉积盆地。

山岩喷发（137~99 Ma，图 8b；卫平生等，2005；钟福平等，2011；汤文豪等，2012；王训练等，2018；陈志鹏等，2019；Hui et al., 2021），地球化学特征指示这些火山岩形成于伸展环境（Hui et al., 2021）。上述伸展断陷盆地大多受控于北东—南西走向正断层或低角度拆离断层或剪切带（Zhou et al., 2012；Zhang et al., 2021b, 2022；Zhao et al., 2025），表明其形成于北西—南东向或北西西—南东东向水平伸展应力作用下（Zhou et al., 2012；Zhang et al., 2018, 2021b, 2022；刘奎等，2024；张进等，2025）。而整个东亚地区，在这一时期整体也处于伸展环境，形成了广泛分布的伸展盆地和变质核杂岩等特征性构造（Davis and Darby, 2010；Wang et al., 2011；Liu et al., 2021；Lin and Wei, 2020；Zhang et al., 2021b, 2022, 2025；Chu et al., 2024；Zhang, 2026）。这些伸展盆地的几何学特征、控盆断裂的断层面滑动矢量数据以及变质核杂岩中的矿物拉伸线理产状，共同指示了该期伸展形成于北西—南东或北西西—南东东向水平伸展应力作用下（Zhang et al., 2021b, 2025；张进等，2025；Zhao et al., 2025）。此次研究在民乐盆地下白垩统的不同层位识别出多处同沉积正断层，控制着下白垩统的沉积（图 3）。这些正断层的原始产状呈现出近南北走向的特征，指示了近东西向的水平伸展（图 3e）。这一控盆构造特征及相应的古应力场方向，与北侧的阿拉善地区及整个东亚地区早白垩世构造特征一致或相近（张进等，2025）。这些证据表明，民乐盆地在早白垩世是一个伸展盆地，其形成受控于近东西向水平伸展作用（图 3）。此外，在民乐盆地西侧的酒泉盆地、花海盆地，以及北侧的潮水盆地和银额盆地中，地震剖面显示盆地内下白垩统中发育大量同沉积正断层，控制了早白垩世地层的沉积（图 9）。断层上盘的下白垩统厚度明显大于下盘（图 9），与民乐盆地野外所见露头尺度的同沉积正断层特征一致。这说明在河西走廊及周边地区，无论从规模还是分布范围上，同沉积正断层在下白垩统中广泛发育，是这一时期的特征性构造。

相关研究已证实河西走廊西端的酒泉盆地在早白垩世为断陷盆地（李奋其，2003；王崇孝等，2005；李明杰，2006；王晓丰等，2008；Zhang et al., 2018），盆地内发育有早白垩世基性火山岩（120~99 Ma；图 8a），其地球化学特征指示了伸展环境（李奋其，2003；汤文豪等，2012）。此次研究认为河西走廊中段的民乐盆地在早白垩世同样形成于伸展环境，并且控制 2 个盆地形成的水平伸展方向也近一致，表明 2 个盆地形成于相同构造背景下。这些证据表明，青藏高原东北缘祁连山以北地区，至少在现今河西走廊地区，在早白垩世处于伸展环境，而非缩短环境。



地震剖面数据来自以下参考文献: 王同和, 1987; 霍永录, 1989; 李明杰, 2006; 冉波等, 2011

a、b、c、d、e—酒泉盆地生长正断层; f、g—花海盆地生长正断层; h、i—银额盆地生长正断层; j—潮水盆地生长正断层

图9 二维地震反射剖面解释所揭示的河西走廊及周边地区早白垩世盆地伸展变形(剖面位置见图8a)

Fig. 9 Extensional deformation of the Early Cretaceous basins in the Hexi Corridor and adjacent areas that indicated by 2D seismic reflection profile interpretation

(a-e) Growth faults in Jiuquan Basin; (f-g) Growth faults in Huahai Basin; (h-i) Growth faults in Yin'e Basin; (j) Growth faults in Chaoshui Basin

4.1.2 晚白垩世构造事件

Wang et al. (2022b) 研究认为民乐盆地在早白垩世沉积之后, 于晚白垩世遭受了缩短变形, 导致下白垩统发生了褶皱。此次研究同样在下白垩统中识别出了与褶皱同期的逆冲断层(图5)。新生代地层与下白垩统之间的角度不整合指示这一缩短变形发生于早白垩世之后、新生代之前(Wang et al., 2022b)。在河西走廊及周边地区, 普遍发育有早、晚白垩世之交的构造运动, 造成了上、下白垩统之间的角度不整合。如民和亭峡峡可

见上白垩统民和组角度不整合于下白垩统河口群之上（甘肃省地质调查院，2021）；民乐盆地以北的潮水盆地内，上白垩统金刚泉组不整合于下白垩统庙沟组之上（Zhang et al., 2017a）。同时，该期挤压还导致早白垩世盆地发生构造反转，盆地内早白垩世的正断层转变为逆冲断层，形成明显的正反转构造（图 9d、9g；霍永录，1989；段宏亮，2004）。此外，河西走廊地区及其北侧的阿拉善地区和南蒙古地区的低温热年代学数据也记录了这一时期的构造热事件（Jolivet, 2017；Zhang et al., 2017a；Song et al., 2018；Zhang et al., 2021b）。该期挤压事件对整个东亚地区均产生了影响，导致早白垩世沉积盆地的构造反转（Suo et al., 2020）以及由低温热年代学数据记录的构造剥露事件（Zhang et al., 2021a；Wang et al., 2024）。同时期的缩短变形广泛发育，在欧亚板块边缘、华南地区以及华北中部，发育强烈的左行转换挤压变形（徐嘉炜等，1985；Darby and Ritts, 2002；Zhang et al., 2020, 2022, 2025），主要受控于北西—南东向或近北向挤压（Zhang et al., 2020, 2025；张进等，2025）。在华北西部的阿拉善地区，发育同时期的挤压褶皱和左行走滑断层，同时受到北东—南西和北西—南东向挤压应力作用（Zhang et al., 2021b, 2022；张进等，2025）。

虽然已有研究认为民乐盆地在晚白垩世主要受到北东—南西向挤压作用，形成了枢纽北西—南东向的褶皱（Wang et al., 2022b）。但此次研究表明，民乐盆地下白垩统形成了北东—南西向和北西—南东向 2 个不同方向的褶皱（图 6a）。虽然不能完全排除两期构造作用叠加的原因，但梨园堡地区敖河向斜所表现出的 type1 型叠加向斜样式（图 6b），进一步证明了该地区下白垩统叠加褶皱作用的发生。这说明民乐盆地在这一时期同时受到了北东—南西向和北西—南东向 2 个方向的水平挤压作用。同时，地层中发育的同时期缩短变形，也分别受控于北西—南东和北东—南西 2 个方向的挤压应力作用（图 5），与北侧阿拉善地区同时期挤压应力场一致（Zhang et al., 2021b, 2022；张进等，2025）。上述证据表明，民乐盆地及周边地区在晚白垩世遭受了北东—南西向和北西—南东向 2 个近直交方向的挤压应力作用，盆内早白垩世沉积地层发生褶皱变形，并形成叠加褶皱（图 6b）。

4.1.3 河西走廊中部晚白垩世热年代学

在河西走廊中部南侧的北祁连造山带前缘，即民乐盆地西侧的榆木山和邻近的金佛寺岩体，已开展了大量低温热年代学数据。其中，金佛寺岩体磷灰石(U-Th)/He 和磷灰石裂变径迹高程-年龄关系表明金佛寺岩体自 9.5 Ma 开始发生强烈挤压隆升（万景林等，2010）；通过对其磷灰石裂变径迹年龄-径迹长度进行模拟，发现金佛寺岩体在晚白垩世（90 Ma）也发生过强烈隆升，之后进入缓慢剥露期，在约 10 Ma 左右再次强烈隆升（Zheng et al., 2010）对万景林等（2010）报导的金佛寺岩体数据重新进行了解释，认为 10 Ma 以来金佛寺岩体向北的水平推覆距离不超过 10 km。而何光玉等（2004）根据玉门油田资料认为北祁连造山带金佛寺段（马营凹陷）向北推复距离可达 48 km。因此，笔者推测这一地区主要推覆期为晚白垩世早期（90 Ma 左右）甚至更早的早白垩世晚期，该期挤压导致酒东盆地晚白垩世至古新世地层的缺失。值得注意的是，紧邻金佛寺的榆木山磷灰石裂变径迹年龄数据也差异性显示了晚白垩世（ $88.7\pm 9.5\sim 62.1\pm 8.2$ Ma）热演化阶段，大部分样品自 90~70 Ma 开始缓慢进入部分退火带（110~60 °C），以恒定剥露速度穿过部分退火带，直至上新世（约 4 Ma）发生快速剥露。部分样品在进入部分退火带后处于构造寂静期（剥露速率约为 0.5 °C/Ma），直至晚中新世（15~5 Ma）快速抬升剥露至地表（Wang et al., 2018）。榆木山的磷灰石裂变径迹冷却历史与紧邻的金佛寺有所区别，可能与地震和大地电磁数据揭示的早白垩世榆木山“飞来峰”构造有关（陈宣华等，2019b）。

在河西走廊中部北侧，主要的山脉为龙首山和合黎山，且均存在晚白垩世热年代学数据报导。龙首山中段磷灰石裂变径迹冷却历史表明，龙首山断裂在晚白垩世（90~60 Ma）陆续快速通过磷灰石裂变径迹部分退火带（110~60 °C），之后长期处于低温（<60 °C）稳定剥蚀状态，直至剥露地表（Zhang et al., 2017a）。民乐盆地和潮水盆地下白垩统古水流指示物源均来自南侧的北祁连山（Vincent and Allen, 1999），晚白垩世以来，龙首山逐渐抬升（Zhang et al., 2017a），使得早白垩世潮水盆地和民乐盆地逐渐分隔，潮水盆地上白垩统金刚泉组开始逐步接收北部阿拉善地区的物源（Vencent and Allen, 1999）。合黎山磷灰石裂变径迹数据表明，该地区大部分样品在早白垩世之前陆续通过部分退火带（110~60 °C），剩余样品则揭示了该地区晚白垩世（ 108 ± 8 Ma 和 83 ± 3 Ma）的构造隆升（An et al., 2018）。

这些低温热年代学数据反映了河西走廊中部在晚白垩世发生的强烈剥露隆升，指示该地区在这一时期经历了构造缩短事件。民乐盆地下白垩统发育的褶皱变形和多种类型的缩短构造，可能就是这一构造事件的直观表

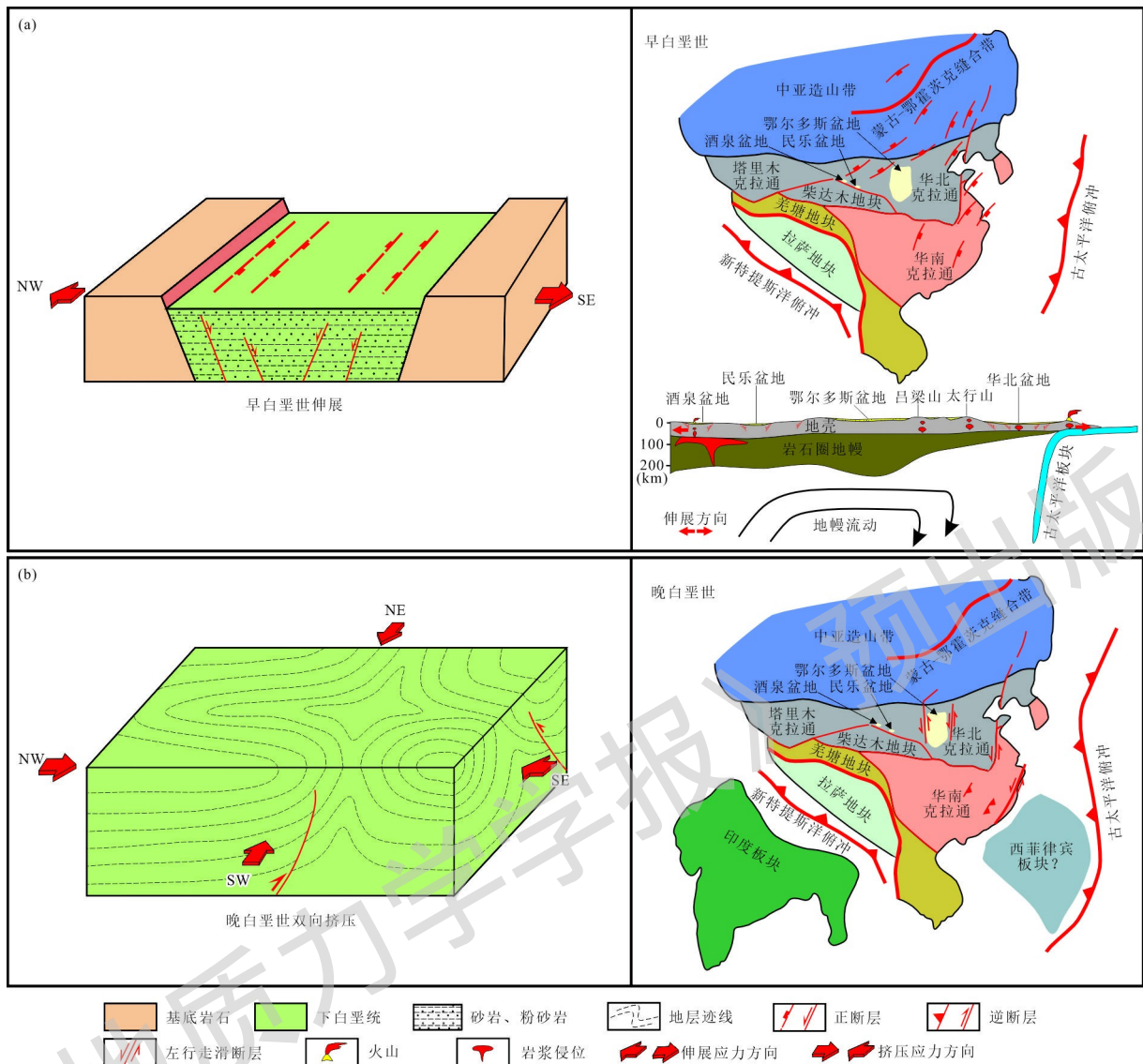
现。

4.2 构造背景

关于东亚地区早白垩世伸展事件的构造机制，相关学者提出了多种解释：①早期加厚地壳的重力垮塌（Wang et al., 2011; Zhang, 2012）；②岩石圈拆沉（Lin and Wei, 2020; Zhu et al., 2018; 董树文等, 2019, 朱日祥和徐义刚, 2019）；③地幔上涌（Zhang et al., 2021c）；④古太平洋板块的俯冲回撤（Liu et al., 2021; Zhu et al., 2017; Zhang et al., 2020）。在民乐盆地周边的阿拉善、北山以及酒泉盆地广泛发育的早白垩世基性火山活动，暗示了这一时期深部地幔的作用（Hui et al., 2021; 韩乐乐等, 2024; 刘奎等, 2024）。Zhang et al. (2020) 根据华北中部山西地区中生代的构造变形研究，提出深部地幔流动对华北早白垩世伸展事件具有重要影响。Hui et al. (2021) 通过对阿拉善地区早白垩世伸展盆地和基性火山岩的相关研究，提出了控制该地区伸展事件的新机制：早白垩世，拉萨地块持续向北俯冲于欧亚板块南缘，古太平洋板块向西俯冲于欧亚板块之下并发生俯冲回撤，在欧亚板块东南缘形成自由边界；同时，俯冲板块的回撤导致地幔物质向东南方向流动，进而引发上部岩石圈向东南方向伸展。这一期伸展向西最远可影响至民乐盆地西侧的北山南部地区（韩乐乐等, 2024; 刘奎等, 2024）。民乐盆地早白垩世的伸展作用受控于近东西向水平伸展应力，与这一时期古太平洋板块回撤以及地幔物质流动方向基本一致，也与同时期东亚地区整体的伸展方向近平行。同时，民乐盆地周边早白垩世盆地内基性火山岩的发育，也暗示了该地区在这一时期的地幔活动（陈志鹏等, 2019; Hui et al., 2021）。因此，由俯冲板块回撤和地幔物质流动共同控制的上部地壳伸展，应为控制民乐盆地早白垩世伸展事件的主要机制（图 10a）。

晚白垩世，在欧亚板块南缘的特提斯构造域和东南缘的太平洋构造域，均发生了重要构造事件。在青藏高原特提斯构造域，拉萨地块持续向北俯冲于羌塘地块之下，造成拉萨地块的抬升以及地壳的缩短（Kapp et al., 2005）；青藏高原东北缘的昆仑山、祁连山等地区在这一时期经历了快速抬升剥露（Yuan et al., 2006; Lin et al., 2021; Liu et al., 2026）。在欧亚板块东南缘，一个未知地块与欧亚板块发生碰撞并沿板块边缘持续向北运动（Zhang et al., 2016; Zhang et al., 2017b; Guo et al., 2018; Suo et al., 2020），产生的挤压作用向北西陆内方向可一直影响至阿拉善地区（Darby and Ritts, 2007; Zhang et al., 2021b, 2022）。这一时期在欧亚板块南缘和东南缘发生的构造事件分别产生了北东—南西向和北西—南东向 2 个近直交的挤压应力场，控制着东亚内陆地区同时期的构造变形。在民乐盆地北侧的阿拉善地区，可明显识别出这 2 个不同方向应力所产生的构造变形（Zhang et al., 2021b, 2022）。此次在民乐盆地下白垩统中同样识别出了轴向北西—南东向和北东—南西向 2 组近直交的褶皱变形，正是北西—南东向和北东—南西向水平缩短作用叠加的结果。同时，地层中发育的同时期缩短变形也分别受控于北西—南东向和北东—南西向 2 个方向的缩短作用。这 2 个应力方向与南侧特提斯构造域和东南侧太平洋构造域同时期构造事件产生的应力场方向基本一致，因此极有可能是这 2 个板块边缘构造事件产生的远程效应在陆内叠加的产物（图 10b）。

综上所述，民乐盆地是一个形成于早白垩世的伸展断陷盆地，随后在晚白垩世经历了多方向挤压改造。该盆地的形成与改造，受控于同时期欧亚板块边缘的构造事件，是板缘应力在大陆内部远程效应的体现。同时，在早白垩世伸展事件中，由板缘构造事件引发的深部地幔物质流动，也为上覆地壳的伸展提供了重要的动力来源。



a—早白垩世民乐伸展断陷盆地模式图及构造背景；b—民乐盆地晚白垩世双向挤压及构造背景

图 10 民乐盆地白垩纪构造演化模式及构造背景

Fig. 10 Cretaceous deformation evolution of the Minle Basin and its tectonic background

(a) Early Cretaceous tectonic model and tectonic background of the extensional deformation in the Minle Basin; (b) Late Cretaceous bidirectional compression and tectonic background of the Minle Basin

5 结论

- (1) 民乐盆地下白垩统中发育大量同沉积正断层，其初始平均走向为近南北向，指示近东西向水平伸展作用，表明民乐盆地为近东西向伸展应力控制的断陷盆地，与同时期东亚地区广泛发育的伸展事件具有一致性。
- (2) 民乐盆地在晚白垩世遭受北西—南东向和北东—南西向双向水平挤压应力的作用，形成了轴向北西—南东向和北东—南西向的叠加褶皱以及多种类型的缩短变形。
- (3) 民乐盆地早白垩世伸展作用主要源于古太平洋板块的回撤及其引发的地幔物质向南东方向的流动，进而带动上覆地壳发生伸展变形。随后在晚白垩世经历的双向水平缩短作用，则是对南侧特提斯构造域和东南

侧太平洋构造域同时期构造挤压事件远程效应在陆内叠加的响应。

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Author's Contributions:

Example: ADINA Alimu and ZHANG Beihang conceived the study, designed the research methodology, conducted the field investigation, performed data analysis, and drafted the manuscript; YANG Yan, ZHAO Heng and JIANG Hai participated in the field investigation and data analysis; LIU Yugang and DING Zhe were involved in the field investigation and manuscript revision. All authors have read and approved the submission and publication of the manuscript.

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